

POLARIZATION-ANGULAR DEPENDENCES OF THE REFLECTION AND TRANSMISSION COEFFICIENTS OF FREE PLATES OF QUASI-ANISOTROPIC ICE IN THE L-RANGE

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Abstract. The polarization-angular dependences of the reflection and transmission coefficients of free plane-parallel ice plates in the region of 0°C, when flat electromagnetic waves of s- and p-polarization fall on them and are recorded by an L-band radiometer 1.41 GHz are studied theoretically. It is shown that the observation of interference effects in ice plates in the L-range is limited to a plate thickness of 4.3 meters. This limitation is due to a radiometer bandwidth of 20 MHz. Ice is considered an isotropic or uniaxial crystal with an anisotropy coefficient of 15%. Three types of directions of the anisotropy axis are considered in the assumptions of the quasi-anisotropic medium model. The value of the Brewster angle is obtained, and it is shown that it does not depend on anisotropy. Reflective and anti-reflective effects are associated with a change in wavelength in the direction perpendicular to the surface of the plate. In turn, this length depends on the angle of incidence, polarization, and direction of the anisotropy axis. The possibility of forming interference regions of transparency for certain thicknesses is shown, the reflection coefficients for both polarizations are significantly reduced, and the transmittances increase. Interference areas of transparency can be formed both at the Brewster angle and at other angles. These effects are associated with interference phenomena and their enhanced effect of anisotropy.

Keywords: reflection, refraction, Fresnel equations, permittivity, millimeter waves, ice, resonance, multilayer plates.

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1. INTRODUCTION

Applications of the general theory of linear electromagnetic waves of anisotropic media [1, 2, 3] require the agreement of general methods based on linear algebra and more traditional approaches to solving diffraction problems. Solutions of many problems of optics [1, 4, 5], acoustics [6], ellipsometry [7], X-ray optics [8, 9] and remote sensing of the environment [10] are based on the use of Fresnel and Fresnel-Airy coefficients. However, in the Fresnel formulas, radiation is assumed to be monochromatic, and the media should be isotropic. The observed interference structures of microwave radiation depend on the bandwidth of the radiometer. Structures are considered theoretically in the case of normal incidence of radiation in

[11]. The class of quasi-anisotropic multilayer plates is considered in [12-14]. This class includes anisotropic media, where s - and p -polarized waves propagate independently. A generalization of Fresnel-Airy formulas to quasi-anisotropic media is considered in [14]. The combination of the methods from [11] and [14] makes it possible to theoretically investigate the effect of anisotropy on the reflection and transmission coefficients of electromagnetic radiation. This paper discusses free ice plates in the L -range.

The Fresnel-Airy coefficients, reflection and transmission coefficients are used in radiometric methods for calculating the brightness temperature of the controlled surface [10]. At frequencies of 1 GHz – 100 GHz, ice, snow, and other types of the earth's surface are studied for a long time [10, 15-22]. In the range above 2 GHz, the behavior of the Fresnel-Airy coefficients of ice plates is complicated by the non-monotonic dependence of the dielectric constant on frequency, as well as by the presence of microcracks [20]. It was reported about the observation of Ginzburg-Pekar waves in this range, near 0°C [24-25]. The complexity of physical phenomena in ice at high frequencies makes the L -band (1-2 GHz) promising for theoretical study. The wavelengths of the L -range are about 20 cm in vacuum and about 10 cm for ice; the Debye formula for the dielectric constant of ice remains valid in it, which simplifies adequate mathematical models. The results of this paper for the L -range should allow to separate the influence of anisotropy from more complex phenomena in other ranges. The presence of mass-produced domestic and foreign L -band radiometers [17, 22, 23] creates the prospect of practical applications of theoretical results.

The practical significance of studying the electric and acoustic anisotropy of ice is substantiated by the significant effect of

anisotropy on the wave processes in sea ice [26]. It is noted that the dielectric anisotropy and anisotropy of the elastic properties of ice may not coincide. In [27], [28], ice is considered as a uniaxial crystal and is characterized by the anisotropy coefficient. For ice-1, the anisotropy of the dielectric constant in the microwave range is approximately 15% [27]. Laboratory experiments to measure the anisotropy of the dielectric constant and specific absorption of arctic drift ice in the microwave range are described in [28], values from 2% to 15% for old ice were obtained, it was noted that the new ice layers are isotropic. Various variants of the direction of the anisotropy axis with respect to the direction of propagation of microwave radiation were considered.

Free plates are of interest for laboratory research, their consideration is a necessary stage in the development of new algorithms. Interference effects in free plates appear more clearly, and the calculation of transmittance is simpler.

In this paper, we restrict ourselves to the consideration of free plates.

The aim of the article is to theoretically consider the influence of ice anisotropy and non-monochromatic radiation on polarization-angle diagrams of reflection and transmission coefficients of L -band electromagnetic waves for free plane-parallel ice plates assuming that the medium is quasi-anisotropic.

2. SELECTION OF COORDINATE AXES AND ASSUMPTIONS OF THE MATHEMATICAL MODEL

In **Fig. 1**, we denote the wave number in parentheses, and the medium number in square brackets. The x -axis is orthogonal to the ice surface. The medium [1] and the medium [3] are air (vacuum). The medium [2] is isotropic or anisotropic ice. The axis of coordinates y is parallel to the projection of the wave vector of the wave in vacuum onto the interface

between the media. The axis of coordinates z is orthogonal to the plane of the figure and is directed to the reader, it lies in the plane of the surface of the plate. The incident wave is indicated by a number 0. The reflected wave has the number 1. The refracted wave in the medium 1 has the number 2, the reflected wave in the medium 2 has the number 3. The wave coming out of the plate has the number 4. When calculating the amplitudes of waves 3 and 4, their multiple reflections are taken into account according to the algorithms [15]. We consider the wave with elliptical polarization as two independently propagating linearly polarized waves with s - and p -polarizations. The electric field of a wave with s -polarization is directed along the z axis. It is convenient to define a wave with p -polarization (like s -polarization), as a wave whose magnetic field is directed along the axis [1]. As the magnitude of the amplitude of a wave with p -polarization, is used. This Such a definition is adequate in media with absorption. In the algorithms used in this work from [15], the angle of refraction is not fundamentally applicable. Instead of them, in accordance with the approach [1], complex wave vectors are considered. Therefore, it is possible to speak of the angle of refraction only formally, as of the real part of the corresponding complex value.

In quasi-anisotropic media, wave vectors (angles of refraction) and wave propagation speeds for s - and p -polarizations should

differ. Further, speaking of x -anisotropy, we will consider a uniaxial crystal, in which the anisotropy axis is directed along the x -axis, orthogonal to the surface of the ice plate. Similarly, we will speak of y - and z - anisotropy as cases when the anisotropy axis is parallel to the ice surface. We assume that there is no spatial dispersion and dispersion of the axes. Then waves with s - and p -polarizations can be identified with the concepts of “ordinary” and “extraordinary” waves adopted in the optics of anisotropic media (see **Appendix**).

Angles of wave propagation are measured relative to the normal to the plate surface. The article deals only with free plates in a vacuum. Further, the terms “angle of incidence”, “angle of observation” are used as equivalent. For a free plate, the angle of incidence is equal to the angle of reflection and equal to the angle of the direction of propagation of the wave leaving the plates, all angles being equal and real. The transmittance and reflection coefficients of the free plates are equal to the square of the modules of the Fresnel-Airy coefficients. $|R_s|^2$, $|T_s|^2$ – reflection and transmission coefficients for waves with s -polarization. $|R_p|^2$ и $|T_p|^2$ – for waves with p -polarization.

In the calculations, the value of the relative complex dielectric constant of ice near $0^\circ C$ is equal to $\epsilon = 3.18 + 0.0007i$ [16, 29, 30]. The value of the imaginary part varied within 15%, real -2% , the changes did not significantly affect the results.

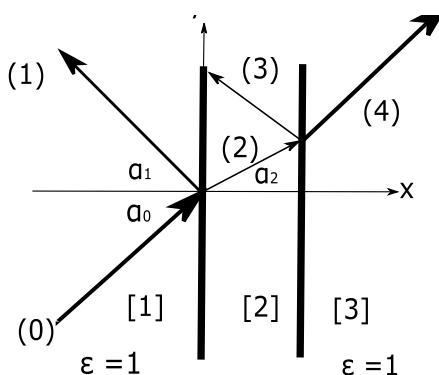


Fig.1. Wave vectors in the free plate.

3. CONDITIONS OF INTERFERENCE EFFECTS FOR FREE PLATES OF ICE AT VARIOUS ANGLES OF INCIDENCE

The wave vector can be decomposed into the ortes of the coordinate axes (e_x , e_y , e_z) and can be represented through the wavelengths in the corresponding directions, noting some conventionality of such a term:

$$\mathbf{k} = \mathbf{k}_x + \mathbf{k}_y + \mathbf{k}_z = \frac{2\pi}{L_x} \mathbf{e}_x + \frac{2\pi}{L_y} \mathbf{e}_y + \frac{2\pi}{L_z} \mathbf{e}_z.$$

Denote by L_x the wavelength in the direction perpendicular to the surface of the plate. For free plates with a multiple of $L_x/2$, the reflection coefficient is significantly reduced. For plates with thickness $L_x/4 + nL_x/2$, (n is an integer), the reflection coefficient increases. (Note that the conditions of extremal reflection are reversed in the case of a monotonic change in the refractive indices in the layers [5]). At a frequency of 14.1 GHz, the wavelength in a vacuum is $L_0 = 21.3$ cm. With normal incidence ($\alpha = 0$), the wavelength of isotropic ice is 11.9 cm. In general

$$L_x = \frac{L_0}{\sqrt{\mu\epsilon - (\sin \alpha)^2}} = \frac{21.3}{\sqrt{3.18 - (\sin \alpha)^2}}.$$

As the angle α increases, the measuring device records the waves with increasing wavelength L_x . Therefore, the conditions of extreme reflection are different at different angles. **Fig. 2** shows the theoretically calculated dependences for isotropic ice on the angle α for L_x , L_y , and the angle of refraction. For isotropic ice, the value of L_x varies from 11.9 cm to 14.4 cm at different angles and leads to interference effects in the plates.

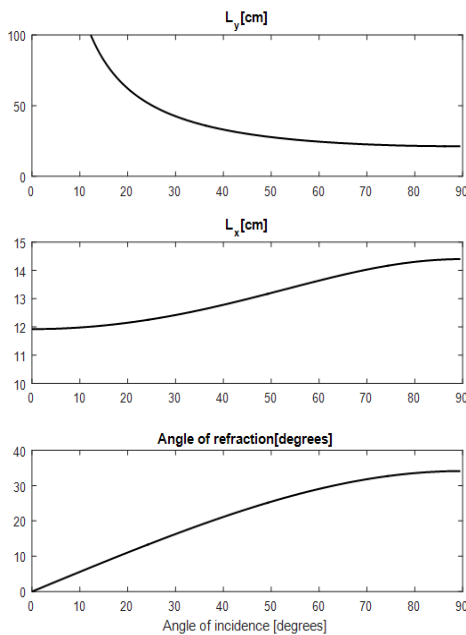


Fig. 2. The dependence of the wavelength L_y , the wavelength L_x and the angle of refraction from the angle of incidence.

4. THE EFFECT OF ANISOTROPY ON THE WAVELENGTH L_x

Appendix shows modifications of Maxwell's equations in quasi-anisotropic media for plane linear harmonic (monochromatic) waves of s - and p -polarization. In the case of x -anisotropy, we denote L_x by L_{xx} . Let $u_x = 0.15$ be the x -anisotropy coefficient. Accordingly, we denote L_{xy} , u_y and L_{xz} , u_z for the cases of y - and z -anisotropies. The relative dielectric constant and other coefficients are written as:

$$\epsilon_{11} = \epsilon(1 + u_x), \epsilon_{22} = \epsilon_{33} = \epsilon; \epsilon_{31} = \epsilon_{32} = \epsilon_{12} = 0; \mu = 1;$$

$$\frac{\omega}{c} = k_0 = \frac{2\pi}{L_0}, k_x = \frac{2\pi}{L_{x,x}}, k_y = k_0 \sin \alpha.$$

x -anisotropy affects only p -polarized waves and does not affect s -polarized waves. The conditions for the existence of solutions for p -waves are converted to:

$$\epsilon^2(1 + u_x) - \epsilon(k_y/k_0)^2 - \epsilon(1 + u_x)(k_x/k_0)^2 = 0.$$

From here we get the wavelength L_{xx} . We represent it in terms of the length L_x for an isotropic medium and the correction for anisotropy proportional to u_x :

$$L_{x,x} = \frac{L_0}{\sqrt{\epsilon - \frac{(\sin \alpha)^2}{1 + u_x}}} \approx L_x - \frac{L_x^3 (\sin \alpha)^2}{2L_0^2} u_x. \tag{1}$$

As L_x should take the values shown in Fig. 2, depending on the angle of incidence. In the case of 15% anisotropy, this correction monotonously grows from zero to 0.6 cm when the angle of incidence varies from zero to 90°.

If the anisotropy axis is directed along the y axis, then the wavelength $L_{x,y}$ is equal to:

$$L_{x,y} = \frac{L_0}{\sqrt{(\epsilon - (\sin \alpha)^2)(1 + u_y)}} \approx L_x - \frac{L_x}{2} u_y. \tag{2}$$

If the anisotropy axis is directed along the z axis, then the anisotropy does not affect the p -polarization. The condition for the existence of a solution for s -polarization takes the form $\epsilon(1$

$+ u_z) - (k_y/k_0)^2 - (k_x/k_0)^2 = 0$. The wavelength $L_{x,z}$ is written as:

$$L_{x,z} = \frac{L_0}{\sqrt{\varepsilon(1+u_z) - (\sin \alpha)^2}} \approx L_x + \frac{\varepsilon L_x^3}{2L_0^2} u_z. \quad (3)$$

In all subsequent graphs, the anisotropic case is represented by solid curves, and the isotropic case is represented by curves consisting of points.

5. ENHANCING THE INFLUENCE OF X-ANISOTROPY BY INTERFERENCE EFFECTS

In accordance with the multiplier $(\sin \alpha)^2$ in formula (1), anisotropy in the x -direction can occur only at sufficiently large angles of incidence and for sufficiently thick plates. For angles greater than 80° , the wavelength in an isotropic medium can be estimated as 14.3 cm. The correction for anisotropy of 15% is 0.6 cm (see section 3), the result is a value close to 15 cm. Therefore, at plate thicknesses of a factor of 7.5 cm, on the curves of the reflection and transmission coefficients, the effects of extreme reflection should be observed for the anisotropic case.

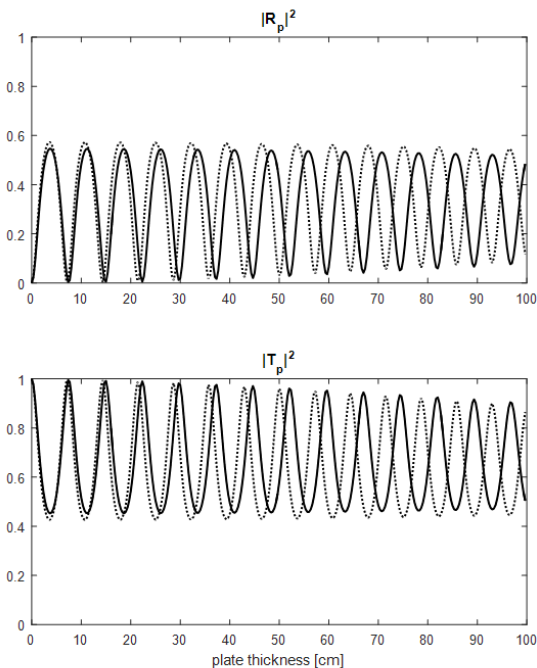


Fig. 3. The dependence of the reflection and transmission coefficients on the thickness of the free plate at an observation angle of 80° .

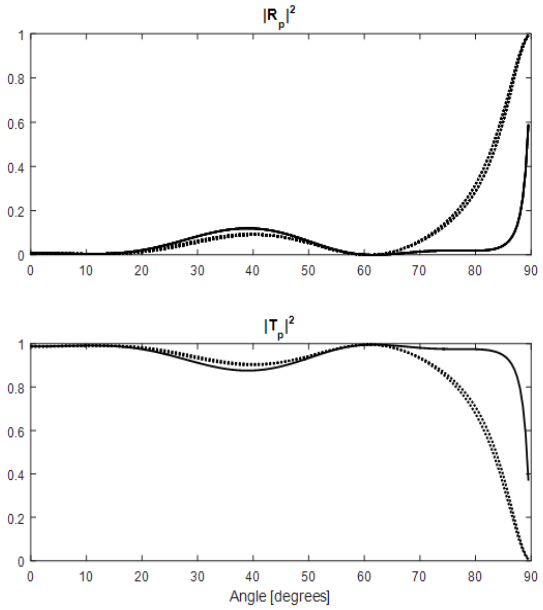


Fig. 4. The reflection and transmission coefficients of ice at a plate thickness of 30 cm.

Fig. 3 shows the reflection and transmission coefficients at ice thicknesses from 0 cm to 100 cm for p -polarized waves incident at an angle of 80° . Due to the difference in the wavelength in the x -direction in an isotropic and anisotropic medium with an ice thickness of 85 cm, the oscillations are in antiphase.

Fig. 4 shows the curves for p -polarization for a thickness of 30 cm. Size 30 cm is a multiple of 6 cm and 7.5 cm. At angles of 0° - 15° , the reflection is small. In addition, for anisotropic plates, low reflection occurs at large angles (60° - 80°).

Recall that the polarization-angle dependences are depicted in the figures in the anisotropic case by solid curves, and in the isotropic case - by curves consisting of points.

6. EXAMPLES OF AREAS OF TRANSPARENCY IN SINGLE-LAYER AND THREE-LAYER PLATES.

Fig. 5 shows the dependence on the angle of incidence of the reflection coefficients and the transmission of a plate 50 cm thick. For isotropic ice, an area of transparency takes place at an angle of 30° . An elliptically polarized wave should pass without reflection and with low

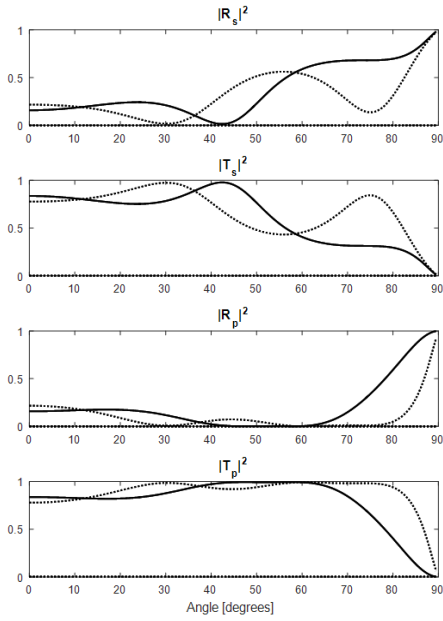


Fig. 5. The reflection and transmission coefficients of ice at a single-layer plate thickness of 50 cm z -anisotropy and y -anisotropy.

losses. The transparency region is destroyed due to the influence of y -anisotropy or z -anisotropy in 15%. The curves of the anisotropic case are shifted relative to the curves of the isotropic case: for s -polarization under the influence of z -anisotropy, or for p -polarization under the influence of y -anisotropy.

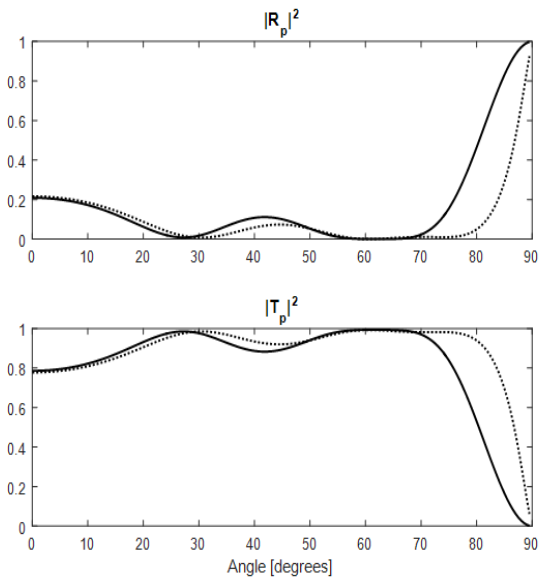


Fig. 6. The reflection and transmission coefficients of ice at a single-layer plate thickness of 50 cm x -anisotropy.

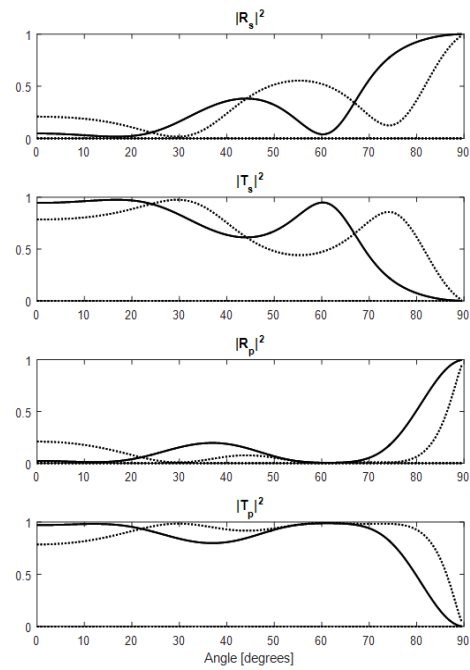


Fig. 7. The reflection and transmission coefficients of ice of a three-layer plate with different anisotropies.

Fig. 6 illustrates the effect of anisotropy in the x direction for a plate with the same thickness of 50 cm as in the previous figure. But the anisotropy axis is 15% orthogonal to the plate surface (x -anisotropy). In the region of 30° the blue line shifts within a few degrees. Transparency region is not destroyed.

Fig. 7 shows the reflection and transmission coefficients for the case when the plate with a total thickness of 50 cm (as in the previous case) is divided into 3 sublayers of the same thickness. In the first sublayer, there is a 15% x -anisotropy, in the second, y -anisotropy, in the third, z -anisotropy. Isotropic and anisotropic ice curves differ at all angles. The region of low reflection of the s -polarization is shifted to the Brewster angle and the transparency region is formed there. As a result of the combined influence of all types of anisotropy, another area of increased transparency is formed at angles of incidence of 0° - 20° .

7. MAXIMUM THICKNESS OF ICE PLATES OF OBSERVATION OF INTERFERENCE STRUCTURES IN THE L-BAND

In Fig. 8, the reflection and transmission coefficient curves for plates with a thickness of 98, 99, 100, 101, 102 cm are shown in the same figure. These thicknesses differ from the average by no more than 2%, but the oscillations of the curves differ significantly and merge into a wide strip. For plates of ice about a meter thick, these oscillations should complicate radiometric measurements.

Fig. 3 shows the variations of the reflection and transmission coefficients at an angle of incidence of 80°. There is a slow decrease in the amplitude of oscillations. It is due not to the absorption of waves, but to interference effects arising from the finite width of the radiometer's passband [11]. The maximum thickness of the plate can be estimated by the formula

$$d = \frac{\lambda f}{2\Delta f} = 423\text{cm},$$

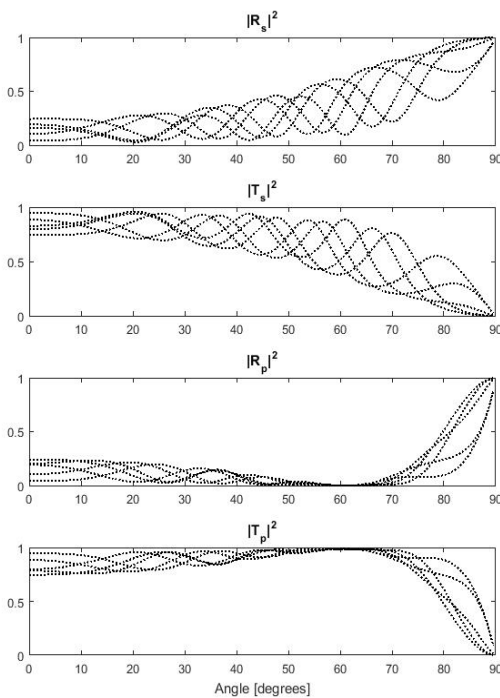


Fig. 8. The reflection and transmission coefficients of isotropic ice with a plate thickness of 98, 99, 100, 101, 102 cm.

λ – wavelength; $f = 14.1\text{GHz}$ – frequency; $\Delta f = 0.02\text{ GHz}$ – radiometer band.

Fig. 9 shows the reflection and transmission coefficients for p -polarized waves with a slightly smaller thickness, 400 cm. Oscillations on curves for anisotropic ice due to interference effects are noticeable. With a larger plate thickness, the amplitudes of interference oscillations decrease, become imperceptible, but exist.

8. CONCLUSION

The polarization-angular dependences of the reflection and transmission coefficients of smooth ice plates for radiometric measurements at a frequency of 1.41 GHz with a bandwidth of 20 MHz are theoretically considered.

When the anisotropy coefficient of ice is 15%, the effect of anisotropy on the reflection and transmission coefficients of infinitely thick plates is insignificant, of the order of a few percent for anisotropy y - and z -, and for x -anisotropy this effect is even less. These estimates are changed when taking into account the thickness of the plates.

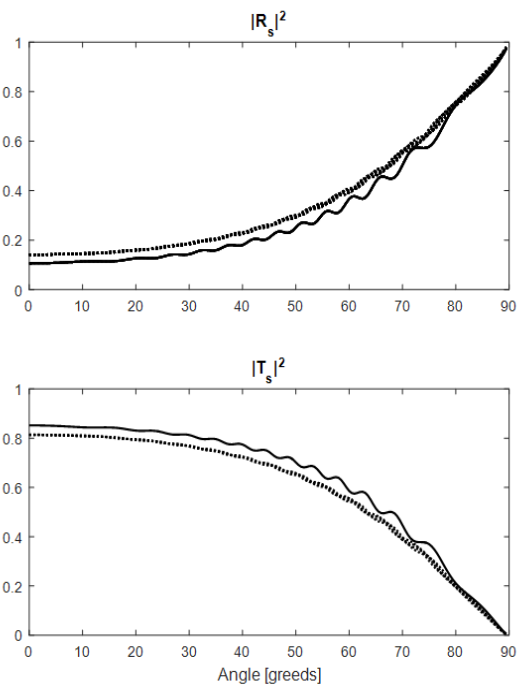


Fig. 9. Reflection and transmission coefficients of s -polarized ice at a thickness of 400 cm.

The wavelength in the direction orthogonal to the surface of the plate inside isotropic ice, at a frequency of 1.41 GHz, is $L_x = 11.9$ cm. This means that the behavior of the reflection and transmission coefficients changes as the thickness of the ice plate changes every $L_x/4 \approx 3$ cm. When the plate thickness changes, there are oscillations of the reflection and transmission coefficients. Between oscillations of the coefficients corresponding to the isotropic and anisotropic case, a phase difference appears that increases with increasing plate thickness, and the effect of relative errors in thickness values on the reflection and transmission coefficients increases. The decrease in the amplitude of oscillations with increasing plate thickness is due to the presence of a limited radiometer frequency band in the frequencies [11].

$$d = \frac{\lambda f}{2\Delta f} = 423 \text{ cm},$$

– maximum thickness of the ice plate, at which interference effects are already invisible.

The knowledge of more accurate values of L_x makes it possible to interpret the phenomena of extremal reflection arising in plane-parallel ice plates less than 423 cm thick, depending on polarization, viewing angle and ice anisotropy. Assuming that anisotropy is small and ice can be considered a uniaxial crystal, in which the anisotropy axis coincides with one of the coordinate axes, formulas for L_x are derived. The versions of formulas in the form of corrections due to anisotropy are given (formulas (1) - (3) in section 3).

The model of quasi-anisotropic media [15] (see also **Appendix**) makes it possible to take into account the anisotropy of the medium, without going beyond the concepts associated with the Fresnel formulas. Interference phenomena enhance the effect of anisotropy. The influence of x -anisotropy becomes significant at angles of more than 70° and ice thicknesses of several tens of centimeters. If the anisotropy axis is parallel to the plane of the plate, then the interference

effects of anisotropy change the reflection and transmission coefficients of the plates in the whole range of angles. The possibility of forming regions of transparency in free plates of ice is theoretically shown. It was verified that the Brewster angle does not depend on the anisotropy of ice. It is shown that transparency regions can be associated with the Brewster angle, but they can also arise at other angles. Anisotropy both destroys areas of transparency, and can contribute to the formation of such areas.

The shorter the wavelength, the smaller the required antenna size. The advantage of the L -range theoretical study is that in this range, it is possible to use simpler mathematical models with dielectric constant values known from the literature, the Debye formula remains valid. While in the range of 3 GHz – 100 GHz, the theoretical explanation of the passage of microwave radiation should be more complex. The obtained numerical estimates of the linear problem can be recalculated for radiometers with shorter wavelengths and used as an approximation for a more detailed study of the properties of ice in other ranges. However, with such a transfer of results, one should take into account the ratio of the radiometer band size to its frequency, the possibility of a more complex behavior of the dielectric constant and its dispersion.

In this paper, only free plates are considered. This fact limits the application of the numerical results obtained mainly by laboratory experiments. For solving the problems of monitoring the earth's surface (and a number of other problems), it would be useful to generalize the algorithms for obtaining transmittance for the case of arbitrary quasi-anisotropic environments. The behavior of the reflection and transmission coefficients for such cases should repeat the same oscillations as in the case of a free plate, with the same wavelengths, but with a different attenuation and phase shift.

The work was carried out within the framework of the state task.

APPENDIX

Permissible components of the dielectric and magnetic permeability tensors for quasi-anisotropic media

Restrictions on the dielectric and magnetic permeability tensors for which TE- and/or TM-waves exist (with the field structure (H_x, H_y, E_z) or $(-E_x, -E_y, H_z)$) were considered in [13–15] and the corresponding generalized Fresnel formulas were obtained. In [15], the conditions for the existence of TE- and TM-waves were proposed to be considered as a definition of a special class of media: quasi-anisotropic media. Quasi-anisotropic media include all isotropic media and some anisotropic media. For multilayer plates of quasi-anisotropic media, algorithms have been constructed that generalize the Fresnel-Airy formulas [15]. The results of this paper were obtained using generalized algorithms and were verified using estimates using the usual Fresnel and Fresnel-Airy formulas. (We also used integration over the radiometer bandwidth and comparing the results with [11]). The concept of quasi-anisotropic media and the ability to solve Fresnel problems for multilayer plates with smooth boundaries from such media expands the range of solvable problems of electromagnetic wave propagation.

Let us write down the permittivity and magnetic permeability tensors admissible for quasi-anisotropic media. The icon \times indicates the permissible components on which the properties of the corresponding waves do not depend.

For s -waves with

$$\hat{\epsilon} = \begin{pmatrix} \times & \times & 0 \\ \times & \times & 0 \\ \times & \times & \epsilon_{33} \end{pmatrix}; \quad \hat{\mu} = \begin{pmatrix} \mu_{11} & \mu_{12} & \times \\ \mu_{21} & \mu_{22} & \times \\ 0 & 0 & \times \end{pmatrix}$$

Maxwell's equations for the amplitudes of monochromatic waves take the form:

$$\begin{pmatrix} \mu_{11} & \mu_{12} & -\frac{c}{\omega}k_y \\ \mu_{21} & \mu_{22} & \frac{c}{\omega}k_x \\ -\frac{c}{\omega}k_y & \frac{c}{\omega}k_x & \epsilon_{33} \end{pmatrix} \begin{pmatrix} H_x \\ H_y \\ E_z \end{pmatrix} = 0.$$

Media of this type may be called s -quasi-anisotropic.

For p -waves at

$$\hat{\epsilon} = \begin{pmatrix} \epsilon_{11} & \epsilon_{12} & \times \\ \epsilon_{21} & \epsilon_{22} & \times \\ 0 & 0 & \times \end{pmatrix}; \quad \hat{\mu} = \begin{pmatrix} \times & \times & 0 \\ \times & \times & 0 \\ \times & \times & \mu_{33} \end{pmatrix}$$

Maxwell's equations for the amplitudes of monochromatic waves take the form:

$$\begin{pmatrix} \epsilon_{11} & \epsilon_{12} & -\frac{c}{\omega}k_y \\ \epsilon_{21} & \epsilon_{22} & \frac{c}{\omega}k_x \\ -\frac{c}{\omega}k_y & \frac{c}{\omega}k_x & \mu_{33} \end{pmatrix} \begin{pmatrix} -E_x \\ -E_y \\ H_z \end{pmatrix} = 0.$$

Media of this type may be called p -quasi-anisotropic.

Maxwell's equations for s - and p -waves are satisfied simultaneously when

$$\hat{\epsilon} = \begin{pmatrix} \epsilon_{11} & \epsilon_{12} & 0 \\ \epsilon_{21} & \epsilon_{22} & 0 \\ 0 & 0 & \epsilon_{33} \end{pmatrix}; \quad \hat{\mu} = \begin{pmatrix} \mu_{11} & \mu_{12} & 0 \\ \mu_{21} & \mu_{22} & 0 \\ 0 & 0 & \mu_{33} \end{pmatrix}.$$

Environments of this type can be called quasi-anisotropic environments in the narrow sense. In media of this type, the s - and p -polarization waves do not interact. Algorithms that generalize the Fresnel and Fresnel-Airy formulas [15] are constructed. In anisotropic media, the wave vectors (wavelengths and angles of refraction) for s - and p -polarizations may differ. In uniaxial crystals, when the direction of the anisotropy axis is parallel or orthogonal to the plate surface, the s -wave and p -polarization waves can be interpreted in terms of the optics of anisotropic media as “ordinary” and “extraordinary” waves. If the anisotropy axis

is directed along the z axis, then a wave with p -polarization is an “ordinary” wave and does not depend on anisotropy; a wave with s -polarization is an “extraordinary” wave and depends on the anisotropy coefficient. If the anisotropy axis lies in the (x, y) plane, the situation is the opposite. A wave with p -polarization is an “extraordinary” wave and depends on anisotropy; a wave with s -polarization is an “ordinary” wave and does not depend on the anisotropy coefficient.

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